



QSR Correspondence

Comment: Radiocarbon deglaciation chronology of the Thunder Bay, Ontario area and implications for ice sheet retreat patterns

Lowell et al. (2009) put forth a hypothesis for the cause of the Moorehead low stand of Lake Agassiz during the Younger Dryas cold event (~12.9–11.5 ka) that differs from the conventional explanation involving eastward routing of the lake to the St. Lawrence River (Broecker et al., 1989). Based on their minimum limiting radiocarbon dated reconstruction of ice-margin retreat from the area of the Lake Agassiz eastern outlet and similar reconstructions for the southern (Fisher and Lowell, 2006; Fisher et al., 2008) and northern (Fisher et al., 2009) outlets, they suggest that Lake Agassiz had no outlets during the Younger Dryas with an outlet opening only after the end of the Younger Dryas. Lowell et al. (2009) thus attribute the Moorehead low stand either to drainage beneath the Laurentide Ice Sheet (LIS) or to enhanced evaporation that had to exceed water inputs from precipitation across the Lake Agassiz drainage basin and meltwater from the LIS.

With regard to their argument for subglacial drainage, Lowell et al. (2009) state that “glaciology identifies possible routes that would drain the basin,” but unfortunately they do not elaborate on how glaciology identifies these routes, where these routes are, or where the water might have gone. To our knowledge, drainage beneath the ice sheet is untenable and unsupported by field evidence: the surface slope of the LIS at this time would have developed a hydraulic gradient that forced water to drain into the lake (i.e., Clarke et al., 2004), a direction that is clearly shown by the numerous eskers in the region (Sado et al., 1994).

With regard to their water-balance argument, the rate of evaporation that would have been required to stabilize Lake Agassiz during this low phase would have been unrealistically high. Based on previous climate modeling and analysis of the LIS hydrology (Kutzbach et al., 1998; Licciardi et al., 1999), a reasonable estimate of the combined volume of net moisture and meltwater discharge into the lake is ~2200 km³ yr⁻¹. This volume of water, distributed over the surface area of the early Moorehead phase of Lake Agassiz (117,000 km²) (Leverington et al., 2000), would have required a minimum evaporation rate of ~18.8 m yr⁻¹ in order to maintain the lake at the Moorehead size. During the late Moorehead phase (185,000 km²) (Leverington et al., 2000), the rate would have been reduced to ~11.9 m yr⁻¹. For perspective, the annual evaporation rate for the Dead Sea is 1.1–1.2 m (Lensky et al., 2005) and 0.4 m for Lake Ontario (Quinn, 1979). Additionally, in a coupled lake-climate simulation of Lake Agassiz (Hostetler et al., 2000), the mean-annual evaporation rate of the lake is 0.2 m. The lake simulation was constrained by geologic evidence that suggests that water temperature of Lake Agassiz may not have exceeded

~5 °C, indicating that substantial solar heating was expended in melting ice and heating cold meltwater. Even if our estimate of the freshwater input to the lake is too high, to achieve a realistic evaporation rate (<1 m yr⁻¹) would require water input rates to the lake to be <190 km³ yr⁻¹, a value well outside the limits of any reasonable water input rate. Based on this analysis, we believe the only explanation for maintaining the lake at the Moorehead low stand is by active drainage through an outlet lower than the southern outlet.

We note that sole use of minimum limiting radiocarbon dates will by default produce nothing but a history of ice-margin retreat. In this vein, Lowell et al. (2009) neglect several publications that suggest a more complex ice-margin retreat history than their simple one developed from minimum limiting radiocarbon dates. Flower et al. (2004) use $\delta^{18}\text{O}$ of seawater ($\delta^{18}\text{O}_{\text{SW}}$) records from the Gulf of Mexico to argue that discharge through the southern outlet (i.e., Mississippi River) abruptly decreased at the start of the Younger Dryas, suggesting abandonment of this outlet several hundred years earlier than the date suggested by Lowell et al. (2009); Fisher and Lowell (2006); Fisher et al. (2008). Furthermore, Carlson et al. (2007) showed that after accounting for the competing effect of temperature on foraminifera $\delta^{18}\text{O}$, there is a significant decrease in $\delta^{18}\text{O}_{\text{SW}}$ in the St. Lawrence estuary at the start of the Younger Dryas, indicating that freshwater was routed from the southern outlet to the eastern outlet. Four additional independent geochemical runoff tracers confirm the routing of Lake Agassiz discharge through the St. Lawrence estuary at the start of the Younger Dryas (Carlson et al., 2007), in agreement with earlier reconstructions (Broecker et al., 1989; Licciardi et al., 1999; Clark et al., 2001).

Finally, Lowell et al. (2009) neglect previous work by Lowell et al. (1999), which dated a readvance of the Lake Superior Lobe across Lake Superior at the end of the Younger Dryas. Lowell et al. (1999) argued that this readvance was contemporaneous along much of the southern LIS extending to eastern Quebec (St. Narcisse Moraine) (LaSalle and Elson, 1975; LaSalle and Shilts, 1993). Given that the majority of the southern LIS readvanced during the Younger Dryas (Lowell et al., 1999), we question the assumption of Lowell et al. (2009) that no readvance occurred just west of Lake Superior, as has been suggested by Teller et al. (2005) and mapped by Clayton (1984). We thus conclude that Lake Agassiz freshwater was routed eastward at the start of the Younger Dryas, in agreement with the original hypothesized forcing of this event (Johnson and McClure, 1976; Rooth, 1982).

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